

IV-2A TWO-DIMENSIONAL FLOW

(Objective: to know how to interpret flow nets)

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(For information only)

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Part IV - 2A TWO-DIMENSIONAL FLOW

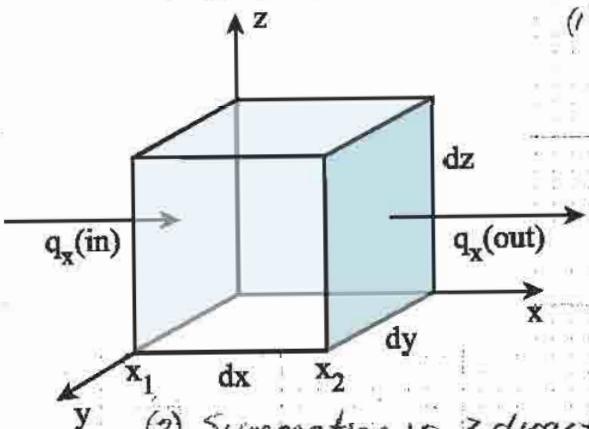
1. BASIC EQUATION

1.1 Assumptions

Darcy's law ($q = kia$) + constant k + incompressible fluid ($S = 100\%$)

1.2 Physical Model for 3-D Flow with changing Volume

- Unit volume with axes coinciding with k axes
 - Partial differentials since $h = f(x, y, z)$



(i) For flow in x direction ($\sigma = dy/dz$; h =total head)

$$g_x(\text{in}) = k_x \left(\frac{\partial h}{\partial x} \right)_{x_1} a \quad g_x(\text{out}) = k_x \left(\frac{\partial h}{\partial x} \right)_{x_2} a$$

$$\Delta g_x = k_x \left[\left(\frac{\partial h}{\partial x} \right)_{x_1} - \left(\frac{\partial h}{\partial x} \right)_{x_2} \right] a$$

Change in gradient over dx

$$= k_x \left(\frac{\partial^2 h}{\partial x^2} dx \right) dy dz$$

Unit volume

$$k_x \frac{\partial^2 h}{\partial x^2} + k_y \frac{\partial^2 h}{\partial y^2} + k_z \frac{\partial^2 h}{\partial z^2} = -\frac{1}{1+\epsilon} \frac{\partial \epsilon}{\partial t}$$

3-D CONSOLIDATION

Net flow into element/unit time
(or out of)

Volumetric change/unit time

1.3 Solution for 2-D Flow & No Volume Change ($\partial e/\partial t = 0$)

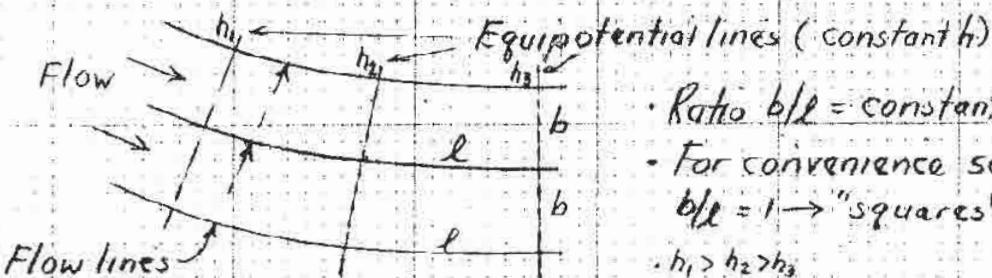
$$(ii) k_x \frac{\partial^2 h}{\partial x^2} + k_3 \frac{\partial^2 h}{\partial z^2} = 0 \rightarrow \frac{\partial^2 h}{\partial x^2} + \frac{\partial^2 h}{\partial z^2} = 0 \quad \text{for isotropic soil (} k_x = k_3 \text{)}$$

Change in gradient per unit distance in x direction

+ " " " is " " " " " 3 " = zero

(2) Laplace Eq. $\nabla^2 h = 0$ (also applicable to heat flow, etc.)

wherein solution \rightarrow series of lines intersecting at 90°



• Ratio $b/l = \text{constant}$

- For convenience select
 $b/f = 1 \rightarrow \text{"squares"}$

$$h_1 > h_2 > h_3$$

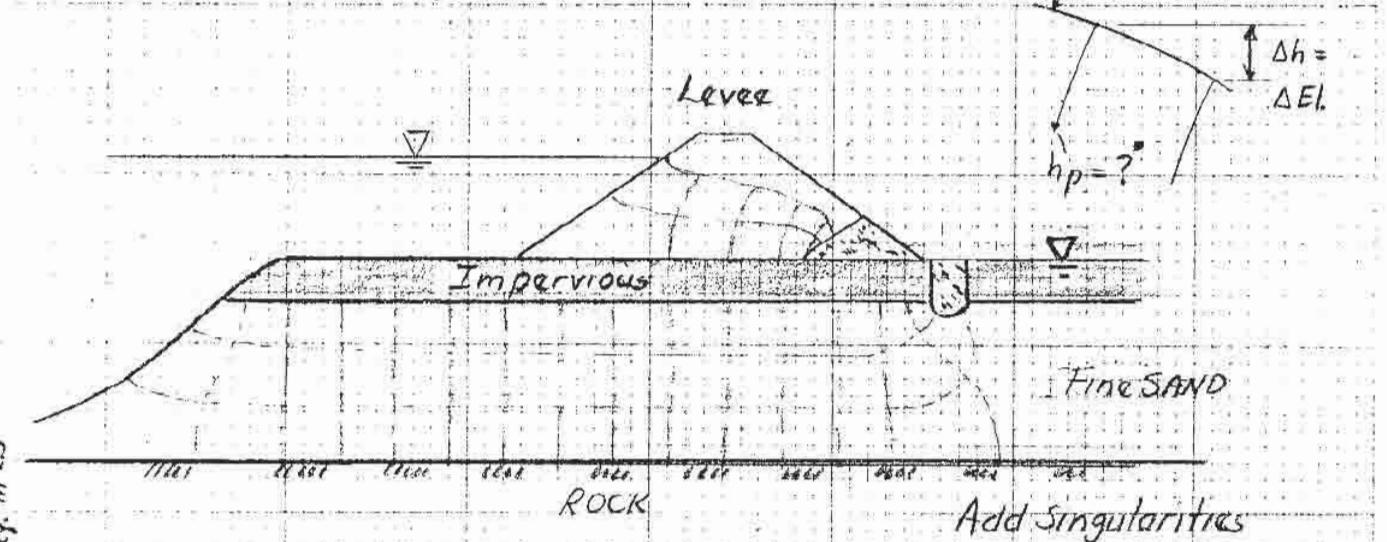
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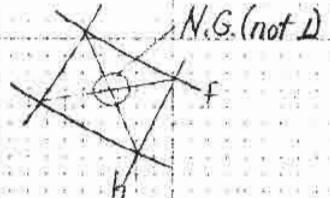
2. DRAWING FLOW NETS

2.1 Problem ($k_h = k_v$)

Add Singularities

2.2 Steps in Drawing Flow Net (If computer program not available)

- (1) Draw problem in ink
- (2) Draw in known equipotential & flow boundary lines
- (3) Sketch in 2 or 3 flow lines (experience helps!)
- (4) Draw corresponding equipotential lines
Check for "squares" & L intersections
- (5) Keep adjusting (and adjusting ...)



2.3 Eq. For Flow Per Unit Length

$$\cdot q \left(\frac{l^3/t}{\ell} \right) = k h \left(\frac{n_f}{n_d} \right) \left(\frac{b}{\ell} \right) = k h (\text{shape factor } \$) \quad \$ \text{ independent of } k \text{ & } h$$

where n_f = no. of flow paths & n_d = no. of equipotential drops

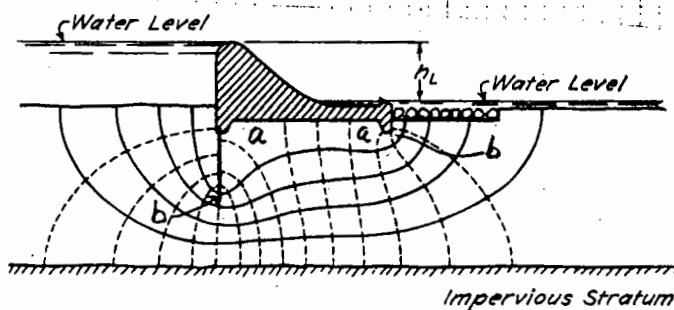
- Even sloppy flow nets (such as above) \rightarrow reasonable \$ and flow estimates given uncertainty in k values.
- BUT sloppy nets can \rightarrow large errors in $i = \frac{\Delta h}{\Delta e}$ at critical locations

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3. MISCELLANEOUS

3.1 Singularities



- (1) When angle between flow & equipotential lines (on side containing the flow)

$$a < 90^\circ \rightarrow l = 0 \text{ (5-sided)}$$

$$b > 90^\circ \rightarrow l = \infty \text{ (3-sided)}$$

- (2) When $l \rightarrow \infty$ near surface, worry about Piping & Quick condition
(3) Identify singularities Fig 2.1

3.2 Non-Homogeneous Soils

- See Sheet A for example that illustrates constant ϕ for varying n_f/n_d & b/e

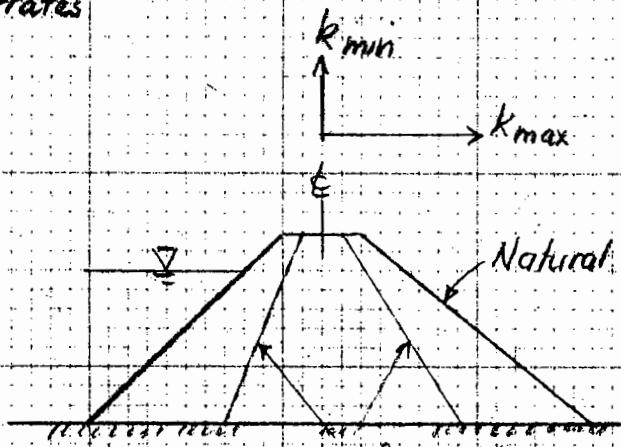
3.3 Anisotropic Soil ($k_x \neq k_y$)

(1) Transformed section:

- Decrease l in direction of k_{\max} by $\sqrt{k_{\min}/k_{\max}}$ OR
- Increase l in direction of k_{\min} by $\sqrt{k_{\max}/k_{\min}}$

(2) Draw flow net $\rightarrow \phi$

$$\text{where } k_e = \sqrt{k_{\min} k_{\max}}$$

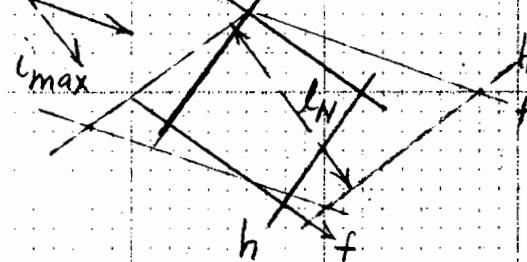


$$q = k_e h \left(\frac{n_f}{n_d} \right) \left(\frac{b}{e} \right)$$

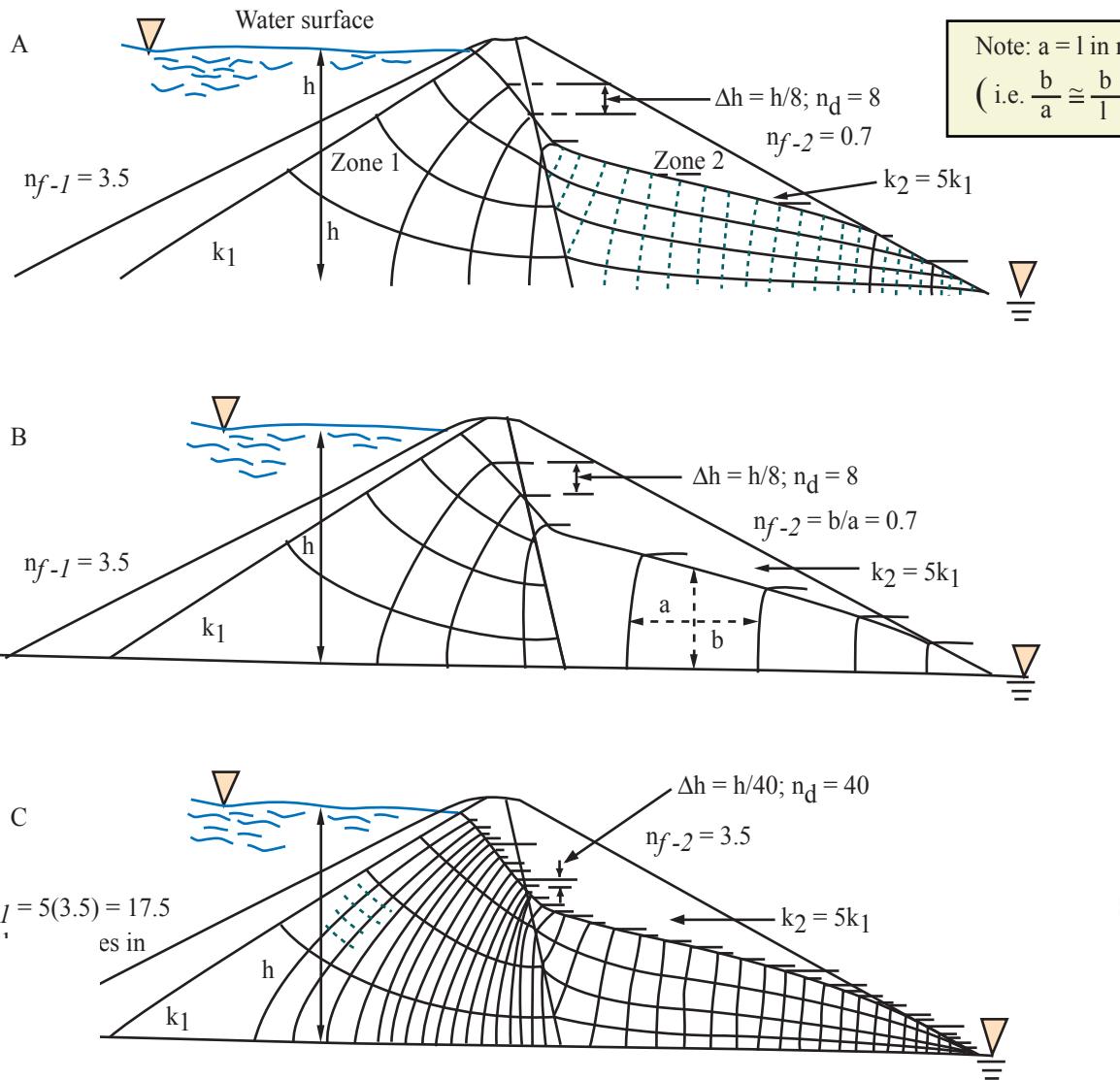
REAL transformed

(3) To obtain i from transformed section ($k_h = \pm k_e$), need \rightarrow natural

Flow



$$i = \frac{\Delta h}{L_N} = \text{shortest distance between equipotential lines}$$



Three forms of one flow net.

$$\text{Fig 4.6(b)} \quad \text{Zone 1: } q_1 = k_1 \frac{h}{8} (3.5)(1) = 0.44 k_1 h$$

$$\text{Zone 2: } q_2 = k_2 \frac{h}{8} (0.7)(1) = 0.088 k_2 h = 0.088 (5k_1) h = 0.44 k_1 h$$

$$\text{Fig 4.6(a)} \quad \text{Zone 2: } q_2 = k_2 \frac{h}{8} (3.5)(\frac{1}{5}) = 0.088 k_2 h = 0.44 k_1 h$$

$$\text{Fig 4.6(c)} \quad \text{Zone 1: } q_1 = k_1 \frac{h}{40} (3.5)(5/1) = 0.44 k_1 h$$

$$\text{Zone 2: } q_2 = k_2 \frac{h}{40} (3.5)(1) = 0.088 k_2 h = 0.44 k_1 h$$

Example of Applying $q = kh(\frac{n_f}{n_d})(\frac{b}{a})$ to Seepage in Non-Homogeneous Soils

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